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Paradox of Peak Flows in a Changing Climate

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13 Almost all observational studies report that extreme precipitation in the U.S. has increased in magnitude over the last several decades 14 153 (Groisman et al. 1999, 2014; Kunkel et al. 2013). Although such studies differ regarding the geographic influence of trends (Mass 164 et al. 2011), the consensus on extreme precipitation (i.e., short-term 17 18 events with a 5% exceedance probability or less) is surprisingly 19 unequivocal. The American Meteorological Society (AMS) 2013 20 State of Knowledge report on extreme precipitation (Kunkel et al. 21 2013) states that "There is strong evidence for a nationally averaged 22 upward trend in the frequency and intensity of extreme precipita-23 tion events ... " and that the in situ measurement network is con-24 sidered "adequate to detect such trends." In a slightly more recent 25 study, Groisman et al. (2013) reported that the very heavy precipi-26 tation rates (in the upper 5% of all records) have increased over approximately two-thirds of the eastern U.S. during the last 27 30 years, whereas the number of days with maximum daily con-28 29 vective available potential energy (CAPE) values exceeding 30 1,500 J/kg has increased \sim 30% in the same period during the 31 spring season. Although the potential causes of this increasing trend may be multifactorial, the most recent AMS report also in-32 33 dicates increasing atmospheric water vapor as a leading causative 34 factor (Kunkel et al. 2013).

Despite the overwhelming consensus on the rising trend of ex-35 treme precipitation rates, the response of peak stream flow is not as 36 37 unequivocal. Although regulation of surface flow, increasing im-38 perviousness, and altering infiltration rates through land cover change are some of the many ways peak flow distribution can 39 be impacted during a stationary precipitation regime, a clear signal 40 41 of the rising trend of extreme precipitation may be expected in peak 42 flow records. However, the paradox of peak flow is that any rising 43 trend in peak flows is much more elusive to observe over statistically significant locations. Vogel et al. (2011) analyzed as many as 44 45 14,000 U.S. streamflow records and found statistically significant 46 increases in flood risk at only approximately 10% of the stations. 47 They also attributed most of this increase to hydrologic changes in land cover (increasing imperviousness, and hence, increased sur-48 face runoff generation) rather than global warming. Villarini et al. 49 50 (2009) analyzed 50 stream flow stations with more than a century-51 long record of flow observations using sophisticated methods for change point detection, trend analysis, and nonstationarity. 52 53 However, they concluded that "it is easier to proclaim the demise 54 of stationarity of flood peaks than to prove it through analyses of 55 annual flood peak data."

Hydrologic extremes are the foundation of most design, operation, and risk management of water management systems that currently serve society. Yet, the current hydrology that traditionally models only the natural laws of physics of a watershed has become increasingly limited in its ability to provide relevant answers for emerging changes that are observed (Vogel 2011). This is because traditional hydrology continues to assume that the extensive replumbing of the natural water network, along with changes to land cover and hemispheric forcing of climate change attributable to extensive human activity, are only an external forcing rather than an integral part of the coupled human–natural system.

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The massive but gradual redistribution of water through artificial reservoirs, numerous irrigation schemes, land cover change, and urbanization since the early 1900s has resulted in a nonnegligible contribution to increased moisture availability and altered atmospheric convergence patterns overland in the U.S. (Puma and Cook 2010; DeAngelis et al. 2010). For example, USGS records (Kenny et al. 2009) indicate an increase in irrigation acreage from 35 million acres (in 1950) to 65 million acres (in 2005). The latter is equivalent to a withdrawal of 144 million acre-ft (or 177 km³) of surface and ground water per year that evaporates directly to the atmosphere [and may be recycled as precipitation (Eltahir and Bras 1996)], and likely balances any increases in peak flow attributable to rising precipitation rates. Similarly, approximately 75,000 artificial reservoirs were built in the U.S. during the last century, with a total capacity almost equaling one year of mean runoff (Graf 1999, 2006; Global Water Systems Project 2008). The cumulative effect of these extensive impoundments has been to triple the average residence time of surface water from 0.1 years (in 1900) to 0.3 years in 2000 (Vorosmarty and Sahagian 2000), an aspect that clearly has not received attention during the assessment of peak flow trends. Similar large-scale alterations have happened to the natural land cover in the U.S., which have modified both hydrologic behavior (water partitioning) and radiative behavior (energy partitioning) in nonnegligible amounts (Pielke et al. 2011).

The explicit consideration of the human replumbing of natural 92 water systems may only explain part of the peak flow paradox. For 93 example, global warming may intensify evaporative fluxes and ini-94 tiate regional drying of soils. This would compensate for increasing 95 precipitation rates through increased abstraction of precipitation 96 and potentially lowering of peak flows. Thus, the hydrologic en-97 gineering community may need to employ a multifactorial ap-98 proach involving, as a fundamental premise, the human impact 99 (feedback) on local-to-regional weather and climate that have been 100 researched for over many decades (Pielke 2001, 2009; Mahmood 6101 et al. 2010), but overlooked in most studies of hydrologic extremes 102 (Hossain et al. 2012). Understanding the causative factors behind 103 the historical evolution of extreme precipitation and peak flow is 104 probably the most urgent priority for current hydrologists to adapt 105 the design and operation of infrastructure. Through an investigation 106 of the combined role of this land-atmosphere feedback owing 107 to artificial redistribution of water, land cover change and climate 108 change, one may postulate that the paradox of peak flows not 109 having responded in sync with rising extreme precipitation may 110 be better understood. Consequently, this approach may allow the 111 hydrologic engineering community to understand how peak flow 112 patterns, which are used in the frequency analysis and design of 113

- 114 many infrastructure, are likely to change in future, given that the 115 replumbing of watersheds will continue. Incidentally, a new term 116 has recently been coined to address such changing nature of hydrol-117 ogy attributable to human impacts. It is designated "hydromorphol-118 ogy" and is defined analogously as the geomorphological approach 119 to hydrology (Vogel 2011).
- 120 There appears to be a need to shift research and education from 121 traditional hydrology to one that addresses hydromorphology for 122 better understanding of this paradox of peak flows. At minimum, 123 this likely entails the consideration of a hydrology curriculum that 124 recognizes the full coupling of the human-natural system with 125 atmospheric feedback (and models), surface-ground-water interac-126 tions, global and regional climate, and interactions with the human 127 management of water resources. Similarly, research would involve 128 the use of atmospheric, hydrologic (surface and ground) models 129 coupled explicitly with water management and redistribution mod-130 eling. Currently, there appears to be no such comprehensive mod-131 eling tool in place that can allow the pursuit of hydromorphology. 132 However, some recent studies, such as that of Biemans et al. (2011), 133 that have aimed to understand the global impact of human manage-134 ment of water by physical systems on local hydrology appears quite promising. If more of such model development work continues 135 136 along this direction, the hydrologic community will one day have 137 the tools required for understanding the hydromorphology of peak 138 flows in the 21st century. In summary, the causes of alteration to 139 peak flow distribution are multifactorial; it may be time for hydro-140 logic engineers to apply a hydromorphological approach rather
- 141 than the purely hydrologic approach of the past.

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